Pore-Scale Heterogeneity and Salinity Impacts on CO₂ Storage in Deep Saline Aquifers: A Microfluidic and Computational Investigation[#]

Rajat Dehury, Jitendra S. Sangwai^{*} Department of Chemical Engineering, Indian Institute of Technology Madras, Chennai, Tamil Nadu 600 036, India. (Corresponding Author: jitendrasangwai@iitm.ac.in)

ABSTRACT

Achieving net-zero emissions and curbing the longterm increase in global warming requires large-scale carbon dioxide (CO₂) sequestration through various techniques. Among the plethora of options for largescale CO₂ storage, geological storage stands out as the most promising, given its expansive storage capacity and minimal environmental impact. Deep saline aquifers, in particular, have emerged as frontrunners due to their vast storage potential and broad global availability. In saline aquifers, various trapping mechanisms, including structural and stratigraphic trapping, residual or capillary trapping, solubility trapping, and mineral trapping, work synergistically to store CO₂. The displacement dynamics of CO₂-brine two-phase flow in pore-scale dictates the structural and residual trapping of CO₂ in subsurface formations. In this regard, a highly heterogeneous porous microchannel was used to investigate the impact of salinity and porous heterogeneity on CO₂-brine displacement dynamics. A computational fluid dynamics (CFD) simulation based on the volume-of-fluid (VOF) method was used to study the immiscible two-phase dynamics under deep reservoir extreme flow temperature and pressure conditions. Microfluidic experiments, conducted under a high-resolution digital microscope with different brine salinities, provide porescale flow visualization and quantification of trapped CO₂. Meanwhile, CFD simulations on the same porous media consider thermophysical property changes for CO₂ and brine under high-pressure and high-temperature conditions, elucidating their impact on CO₂ trapping or storage. The observed unstable finger-like displacement pattern in the microchannel, attributed to very low capillary number (Ca) and viscosity ratio (M) values, results in CO₂ channeling with limited water displacement, thus restricting CO₂ saturation at around 50 %. Along with viscous fingering, small pore-throats pronounce the snap-off effects at the micro-scale, which increase the residual trapping of CO_2 in the porous media. An increase in brine salinity increases viscosity, interfacial tension, and contact angle, leading to a more vertical sweep of water and increasing the CO_2 storage capacity of a saline aquifer. Under reservoir conditions, the comparatively higher density of CO_2 and increase in contact angle resulting from the adsorption of CO_2 increases CO_2 trapping in porous channels. This combined approach of microfluidic experiments and CFD simulations provided valuable visualization and insights into CO_2 -brine dynamics at the pore-scale, which would contribute significantly towards efficient, enhanced, and secure storage of CO_2 .

Keywords: capillary number, CO₂ geo-sequestration, micromodel, porous media, saline aquifers, snap-off effect

NOMENCLATURE

Abbreviations	
CCUS	Carbon capture, utilization, and sequestration
CFD	Computational fluid dynamics
IEA	International Energy Agency
IFT	Interfacial tension
Gt	Gigatons
VOF	Volume of fluid
Symbols	
Са	Capillary number
E	Total mass-averaged energy
\vec{F}_{csf}	Interfacial force term
h	Enthalpy
Ĵ	Total diffusive flux
k	Interface curvature
K _{eff}	Effective conductivity
Μ	Viscosity ratio
ñ	Unit surface normal
Ρ	Pressure

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t	time
Т	Temperature
$ec{v}$	Velocity
α	Volume fraction
μ	Viscosity
Σ	Interfacial tension
θ_w	Contact angle
Ρ	Density
$\overline{\overline{ au}}$	Viscous stress tensor

1. INTRODUCTION

According to the International Energy Agency (IEA), the world's energy supply is projected to increase by 21% by 2050 compared to 2022 under currently stated policy scenarios due to several factors, including economic growth, population growth, and technological advances [1]. As of 2023, fossil fuels still dominate the world energy sector, contributing to over 80 % of the world's total energy demand [2]. Annual carbon dioxide (CO₂) emissions also increased up to around 36.3 Gt/yr (gigatons per year) in 2021, with an average CO₂ level of 416 ppm in the atmosphere [3, 4]. Therefore, immediate actions must be taken to meet the Paris Agreement goal of limiting global temperature by 1.5°C and achieving net-zero scenarios. Various ways for transitioning to netzero economies involve transitioning to renewable energy sources, improving the energy efficiency of different energy extensive processes, implementing large-scale carbon capture, utilization, and sequestration (CCUS) methods, improving policy and regulatory processes for climate change, and promoting international cooperation for clean technologies. Considering the dependence of the energy sector on fossil fuels, it is evident that other options like CCUS should be regarded rather than focusing only on renewable energy sources. After capturing carbon dioxide, it must be sequestered in underground geological formations or deep ocean waters.

Considering safety and environmental concerns for marine life, large capacities of underground porous media, monitoring accessibility, and public acceptance, geological storage seems to be the better option at this time [5]. Underground formations suitable for CO₂ sequestration can be either porous media like depleted oil and gas reservoirs and aquifers or non-porous media like host rocks, salt caverns, and abandoned mines. In porous media, gas is stored in interconnected pores or voids previously saturated with saline water or hydrocarbon fluids [6]. Along with depleted reservoirs, saline aquifers are also being explored by the scientific community for large-scale CO₂ storage due to their broader availability, versatility with depth and size, reduced leakage and environmental impact (due to mixing with the oil), and potential dual-purpose use as enhanced geothermal systems [7, 8]. There are four primary trapping mechanisms that aid in the safe storage of CO_2 in underground porous media such as structural/stratigraphical, residual, solubility, and mineral trapping [9]. All four trapping mechanisms contribute towards secure CO_2 sequestration in the case of subsurface saline aquifers.

An accurate assessment of CO₂-brine displacement dynamics in the pore scale determines the structural, solubility, and residual trapping of CO₂. However, limited pore-scale studies have discussed gas displacement dynamics and the corresponding trapping capacity of porous media in terms of both qualitative and quantitative analysis. Moreover, gas-liquid two-phase flow in porous media gets highly complex due to unstable displacement patterns, which are not well investigated. Experimental studies involving microfluidic investigation, 3D X-ray microtomography technique, numerical simulations like lattice Boltzmann method, volume of fluid (VOF) method, and pore-network modeling approach have been utilized by researchers for pore-scale characterization of CO₂ trapping and storage capacity estimation [10-17]. Zhang et al. conducted a investigation microfluidic and compared the displacement efficiency of nitrogen (N₂), hydrocarbon gas, and CO₂ for miscible and immiscible displacement cases. The discussion focuses on geological CO₂ utilization and storage based on enhanced oil recovery and concludes that CO₂ is better for oil production, especially in deep reservoirs with high pressure [10]. Al-Zaidi et al. and Moghadasi et al. used the X-ray microtomography technique to understand CO₂-brine displacement dynamics in core samples in terms of pressure behavior. The unstable displacement dynamics are highlighted in these research articles [11], [14]. Basirat et al. used a numerical modeling technique in a 2D microscopic heterogeneous model with the phase field method for interface calculations between CO₂ and brine for deep investigation of underground carbon sequestration. The studies show that with a decrease in contact angle, the wetting phase (brine) saturation and normalized interfacial increases [12].

The microfluidic investigation provides a direct visualization during two-phase displacement experiments, allowing for the understanding of unstable displacement patterns like viscous and capillary fingering. In this study, both experimental and simulation techniques have been used to better understand the displacement dynamics between CO_2 and brine. The same microfluidic model was used for both experimental and simulation studies, which mimics subsurface porous media. The study aims to quantify the structural and residual trapping of CO_2 and discuss the impact of pore heterogeneity and salinity on storage capacity.

2. MATERIALS AND METHODS

2.1 Numerical Methodology

To simulate CO₂-brine two-phase pore-scale flow, a computational fluid dynamics (CFD) simulation coupling with volume of fluid (VOF) method was used. The VOF model was used to track the interface between the two phases during the simulation. The VOF method introduced a new variable called volume fraction (α), and the corresponding continuity equation for α was solved for the interface tracking. Along with mass, momentum, and energy conservation principles, interfacial forces like viscous and capillary forces were introduced in the model to capture the unstable fingering phenomena at the pore scale. All the governing equations used in this simulation are shown in Table 1.

2.2 Porous Geometry and Experimental Setup

The porous media used for this study has a dimension of 20 mm \times 10 mm \times 20 μ m, with 2.5D permeability and 0.57 porosity. For simulation purposes, a two-dimensional (2D) geometry was considered due to the small depth of the three-dimensional micromodel. The average pore size and throat size radii of the physical rock-type microchannel are 46 and 56 μ m, respectively. The microfluidic device, as shown in Fig. 1a, has intricate and highly tortuous flow paths and mimics underground porous media. The microfluidic experiments to investigate CO₂-brine displacement dynamics were conducted under a digital microscope with a camera to capture the flow dynamics. Fig. 1b illustrates a simplified experimental setup with a syringe pump for CO₂/brine injection, a microchannel for flow experiments, a digital microscope and camera setup for flow visualization, and a computer for flow investigation. Five different salinities, 0, 10, 30, 60, and 80 parts per thousand (ppt) of water, were used to study the effect of salinity on twophase displacement dynamics. Various salinity values were obtained from seawater (with approximately 30

Table 1 The governing equations used to simulate CO₂-brine two-phase immiscible flow in a heterogeneous porous media under subsurface reservoir conditions.

Equation Names	Mathematical Representation		References
Volume fraction tracking	$\frac{\partial}{\partial t} (\rho_1 \alpha_1) + \nabla (\rho_1 \alpha_1 \vec{v}) = 0$	(1)	[18, 19]
Navier-Stokes equations	$\frac{\partial}{\partial t}(\rho \vec{v}) + \nabla (\rho \vec{v} \vec{v}) = -\nabla p + \nabla (\mu (\nabla \vec{v} + \nabla \vec{v}^T)) + \vec{F}_{csf}$	(2)	[19, 20]
Surface force and wall adhesion term	$\vec{F}_{csf} = \sigma \frac{\rho k \nabla \alpha_1}{\frac{1}{2} (\alpha_1 + \alpha_2)}$	(3)	
	$2^{(\mathcal{V}_1 + \mathcal{V}_2)}$ $k = -(\mathcal{V}.\hat{n})$	(4)	[21, 22]
	$\hat{n} = \hat{n}_w cos\theta_w + \hat{t}_w sin\theta_w$	(5)	
Energy equation	$\begin{aligned} \frac{\partial}{\partial t}(\rho E) + \nabla &\cdot \left(\vec{v}(\rho E + p)\right) \\ &= \nabla \cdot \left(K_{eff}\nabla T - \Sigma\Sigma \ h_{j,1}\vec{J}_{j,1} + (\bar{\tau} \cdot \vec{v})\right) \end{aligned}$	(6)	[19, 20]
Properties at the interface	$\rho = \alpha_1 \rho_1 + (1 - \alpha_1) \rho_2$	(7)	[40]
	$\mu = \alpha_1 \mu_1 + (1 - \alpha_1) \mu_2$	(8)	[18]

ppt), either by diluting with distilled water or by evaporating the seawater. The water phase was dyed using a blue color dye to differentiate the two phases easily during the flow displacement experiments. Initially, the micromodel was completely saturated with brine. Following the brine saturation, it was left undisturbed for 12–24 hours to achieve a more accurate approximation of subsurface aquifers. CO₂ was injected from one side of the microchip to observe the displacement dynamics between the two phases. For each salinity case, experiments were repeated 2-3 times, and average values of flow saturation were reported. After each experiment, the chip was cleaned with ethanol, alkali, and toluene to remove any impurities and previous saline waters.





Fig. 1 (a) Heterogeneous porous media used for the two-phase flow investigation; (b) Experimental setup for microfluidic experiments.

2.3 Pore-scale Studies

All the microfluidic studies were carried out in atmospheric conditions, while numerical simulations were used to investigate under atmospheric and aquifer conditions (P = 15 MPa and T = 323 K). To understand the flow displacement regimes, all the simulation and experimental cases were plotted on a capillary number (*Ca*) vs. viscosity ratio (*M*) phase diagram, as shown in Fig. 2. *Ca* is defined as the ratio of the viscous force of

injected fluid to capillary force $(Ca = \frac{\mu_2 u}{\sigma})$ and M is the ratio of viscous force between displacing and displaced fluid $(M = \frac{\mu_2}{\mu_1})$. All the considered cases lie between the viscous and capillary fingering regions, which allows to study the influence of viscous and capillary forces on displacement dynamics at the pore scale. The displacement regimes are divided following the works of Lenormand et al. and Zhang et al. [23, 24].



Fig.2 Log (Ca) vs. Log (M) phase diagram illustrating cases considered for this study and three different flow regimes following the work of references [23, 24].

Before carrying out the simulations discussed in this study, the CFD model was checked for mesh independence and validated by comparing it with microfluidic experiments. A detailed discussion on validation is available in our previous literature [20]. Along with atmospheric pressure, simulations were performed under high-pressure and high-temperature conditions observed in subsurface reservoirs. CO₂ becomes supercritical under these conditions, and its properties are calculated following the work of Span and Wagner [25]. The drastic change in CO₂ density and viscosity observed due to high pressure and temperature cases can be visualized in Fig. 3a and b. Brine and interfacial properties at these conditions were determined following previous literature [20, 26, 27].

3. RESULTS AND DISCUSSIONS

3.1 Experimental Observations

All the microfluidic experiments were conducted under atmospheric conditions to observe the brine displacement patterns and corresponding flow dynamics. The microfluidic device consists of multiple high-permeable, mid-permeable, low-permeable, and dead-end regions to mimic a physical rock found in subsurface saline reservoirs on a guasi-2D scale. A typical observation for а microfluidic experiment is



Fig. 3 Variation of CO₂ properties at high-pressure and high-temperature geological aquifer conditions (a) Density; (b) Viscosity.

demonstrated in Fig. 4a. Initially, a clean microchip was saturated with brine, and CO₂ was flooded. In atmospheric conditions, the interphase is clearly differentiable. indicating an immiscible phase displacement. To understand the physical displacement patterns at the pore scale, 8 locations on the microchip were identified with different permeabilities and at various distances from the injection side. Fig. 4b-d illustrates a few critical observations from the microfluidic experiments. CO₂ initially prefers a highpermeable flow path, creating multiple finger-like structures and channeling through these high-permeable channels (Fig. 4b and c). Viscous fingers developed due to the high viscosity contrast (M =0.01) between brine and CO₂, which can be observed throughout the micromodel. Fingers developed in the high-permeable

path grow faster and create a positive feedback mechanism, allowing subsequent injected CO₂ to follow the same paths [20, 28]. Fingering patterns of mid and low-permeable flow paths only develop after the invasion of high-permeable micro channels (Fig. 4b). Fig. 4c illustrates the CO₂-channeling phenomena and poor sweep efficiency of dead-end regions. This CO₂ channeling through high-permeable channels limits the stratigraphical trapping of CO₂. The average CO₂ saturation observed for 10 ppt salinity brine invasion is 0.495. Fig. 4d shows the snapped-off CO₂ bubbles generated as the non-wetting phase separates from the continuous phase and moves separately.



Fig. 4 Microfluidic experiments for the study of CO₂brine (10 ppt) displacement dynamics. (a) A typical CO₂ flooding observation; (b) Different degrees of viscous fingering; (c) CO₂ channeling (d) CO₂ snap-off events.

3.2 Observations from CO₂ Pore-scale Simulations

CO2-brine two-phase numerical simulations were carried out both at atmospheric and reservoir conditions. Under reservoir conditions, due to the supercritical nature of CO₂, breakthrough time and saturation at breakthrough increased significantly. The displacement patterns obtained from CFD simulations highlight the unstable nature of the brine invasion on a pore scale. For 10 ppt salinity values at the atmospheric conditions, Ca and M lie in the intermediate region between capillary and viscous fingering, and the corresponding displacement patterns are demonstrated in Fig. 5. The displacement patterns show unstable fingering patterns and snapped-off CO₂ bubbles. Snapped-off bubbles are generated when a gas bubble elongates while traveling through a narrow porous channel and gets disconnected from the continuous phase. A comprehensive understanding of the snap-off effect can help better

estimate CO₂ residual trapping inside a porous media. Higher degrees of snap-off events increase the residual trapping of CO₂ and may also increase the solubility trapping as the interfacial area between CO₂-brine increases with each snap-off event. Along with the channel geometry, the degree of snap-off events depends upon the flow rate of the injection, interfacial tension (IFT), and contact angle between rock-brine-CO₂. This study observed higher snap-off events for lowsalinity brine, which can be attributed to their low contact angle and low IFT. The CO₂ breakthrough time for 10 ppt saline water is 0.739 s, and the CO2 saturation at breakthrough is 0.367. Microfluidic experiments generally estimate higher CO₂ saturation as these studies calculate saturation values after a steady state. However, in simulation cases, saturation values were obtained immediately after the breakthrough of CO₂ at the outlet. Comparing CO₂ saturation at breakthrough can be a better practice to have safe and efficient storage.



time stages of CO₂-brine for 10 ppt salinity. Breakthrough time = 0.739 s and saturation at breakthrough = 0.367.

3.3 Impact of Salinity on CO₂ Trapping

Brine with five different salinity values, 0, 10, 30, 60, and 80 ppt, were used to investigate the impact of brine salinity on CO_2 trapping by stratigraphical and residual mechanisms and CO_2 -brine displacement patterns. Simulations were conducted at atmospheric and highpressure (P= 15 MPa) and high-temperature (T= 323 K) conditions. Fig. 6 shows a comparative plot of breakthrough time and saturation at breakthrough for different simulation cases. As illustrated with increased brine salinity, both breakthrough time and saturation increase. The trend is similar for both atmospheric and reservoir conditions. After 60 ppt salinity, CO_2 saturation remains constant, suggesting the influence of high viscous force dominating interfacial force compared to other salinity values. Even though the trend is similar for both atmospheric and aquifer conditions, the breakthrough time and CO_2 saturation increased by an average of 37.5 % and 29.1 % for underground aquifer conditions. The results show higher CO_2 storage capacity of high-temperature and high-pressure reservoirs due to the supercritical nature of CO_2 . High density and high viscosity at reservoir conditions (Fig. 3) increased the time taken by CO_2 to reach the outlet of the micromodel.



Fig. 6 CO₂ breakthrough time and CO₂ saturation at breakthrough for different salinity values both at atmospheric and geological reservoir conditions.

 CO_2 -brine displacement dynamics highly depend upon the interfacial forces, IFT, and contact angle. The variation of these two parameters determines the CO_2 storage capacity of the micromodel with varying salinity. Fig. 7a and b show the trend of IFT and contact angle for different salinity values at atmospheric and reservoir conditions. Both interfacial tension and contact angle increase with increased brine salinity for atmospheric and reservoir conditions. While in reservoir conditions, IFT decreases drastically, and the contact angle increases, making the system less water-wet. As a result, high-salinity brine is expected to displace easily, allowing more CO_2 storage capacity. Similar observations are obtained from the simulation results under extreme reservoir conditions.

4. CONCLUSION

CO₂ sequestrations will play a crucial role in achieving a net-zero economy. Subsurface saline aquifers have been explored as a potential site for large-scale, safe, and efficient carbon storage. In this context, this study explores the pore-scale flow dynamics of CO₂brine, which helps understand CO₂ storage capacity and



Fig. 7 (a) Interfacial tension and (b) Contact angle for different salinity values at atmospheric and reservoir conditions.

leakage status for a porous media filled with saline water. Five different salinity values, 0, 10, 30, 60, and 80 ppt, were considered to evaluate the impact of porescale heterogeneity and salinity on CO₂-brine two-phase flow in a heterogeneous porous microchannel. Microfluidic experiments were performed under atmospheric conditions, and numerical CFD simulations were conducted at atmospheric and high-pressure hightemperature conditions. Microfluidic experiments show the CO₂-brine displacement dynamics at various sections of the microchip. Degrees of viscous fingering, CO₂ channeling, and snap-off events characterize the CO₂brine displacement dynamics at the pore scale. For reservoir conditions, CO₂ storage capacity increased by 29.1 % due to the high density and viscosity of supercritical CO₂. Storage capacity increases with the increase in salinity as the system becomes water-wet with higher salinity. Saline aquifers with a 30-60 ppt salinity can be used to get maximum storage capacity with a balance of viscous and capillary forces. Higher salinity also shifts the displacement dynamics slightly towards the capillary fingering regime. This combined approach of microfluidic experiments and CFD simulations provides a visualization of CO₂-brine two-phase displacement dynamics and helps predict the storage capacity of porous media, which would contribute towards a successful large-scale CO₂ sequestration.

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