Pore-Scale Heterogeneity and Salinity Impacts on CO² Storage in Deep Saline Aquifers: A Microfluidic and Computational Investigation

Rajat Dehury, Jitendra S. Sangwai* Department of Chemical Engineering, Indian Institute of Technology Madras, Chennai, Tamil Nadu 600 036, India. (Corresponding Author: jitendrasangwai@iitm.ac.in)

ABSTRACT

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Achieving net-zero emissions and curbing the longterm increase in global warming requires large-scale carbon dioxide $(CO₂)$ sequestration through various techniques. Among the plethora of options for largescale $CO₂$ storage, geological storage stands out as the most promising, given its expansive storage capacity and minimal environmental impact. Deep saline aquifers, in particular, have emerged as frontrunners due to their vast storage potential and broad global availability. In saline aquifers, various trapping mechanisms, including structural and stratigraphic trapping, residual or capillary trapping, solubility trapping, and mineral trapping, work synergistically to store $CO₂$. The displacement dynamics of CO2-brine two-phase flow in pore-scale dictates the structural and residual trapping of $CO₂$ in subsurface formations. In this regard, a highly heterogeneous porous microchannel was used to investigate the impact of salinity and porous heterogeneity on $CO₂$ -brine displacement dynamics. A computational fluid dynamics (CFD) simulation based on the volume-of-fluid (VOF) method was used to study the immiscible two-phase flow dynamics under deep reservoir extreme temperature and pressure conditions. Microfluidic experiments, conducted under a high-resolution digital microscope with different brine salinities, provide porescale flow visualization and quantification of trapped CO2. Meanwhile, CFD simulations on the same porous media consider thermophysical property changes for $CO₂$ and brine under high-pressure and high-temperature conditions, elucidating their impact on $CO₂$ trapping or storage. The observed unstable finger-like displacement pattern in the microchannel, attributed to very low capillary number (Ca) and viscosity ratio (M) values, results in $CO₂$ channeling with limited water displacement, thus restricting $CO₂$ saturation at around 50 %. Along with viscous fingering, small pore-throats pronounce the snap-off effects at the micro-scale, which increase the residual trapping of $CO₂$ in the porous media. An increase in brine salinity increases viscosity, interfacial tension, and contact angle, leading to a more vertical sweep of water and increasing the $CO₂$ storage capacity of a saline aquifer. Under reservoir conditions, the comparatively higher density of $CO₂$ and increase in contact angle resulting from the adsorption of $CO₂$ increases $CO₂$ trapping in porous channels. This combined approach of microfluidic experiments and CFD simulations provided valuable visualization and insights into $CO₂$ -brine dynamics at the pore-scale, which would contribute significantly towards efficient, enhanced, and secure storage of CO₂.

Keywords: capillary number, CO₂ geo-sequestration, micromodel, porous media, saline aquifers, snap-off effect

NOMENCLATURE

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1. INTRODUCTION

According to the International Energy Agency (IEA), the world's energy supply is projected to increase by 21% by 2050 compared to 2022 under currently stated policy scenarios due to several factors, including economic growth, population growth, and technological advances [1]. As of 2023, fossil fuels still dominate the world energy sector, contributing to over 80 % of the world's total energy demand $[2]$. Annual carbon dioxide $(CO₂)$ emissions also increased up to around 36.3 Gt/yr (gigatons per year) in 2021, with an average $CO₂$ level of 416 ppm in the atmosphere [3, 4]. Therefore, immediate actions must be taken to meet the Paris Agreement goal of limiting global temperature by 1.5°C and achieving net-zero scenarios. Various ways for transitioning to netzero economies involve transitioning to renewable energy sources, improving the energy efficiency of different energy extensive processes, implementing large-scale carbon capture, utilization, and sequestration (CCUS) methods, improving policy and regulatory processes for climate change, and promoting international cooperation for clean technologies. Considering the dependence of the energy sector on fossil fuels, it is evident that other options like CCUS should be regarded rather than focusing only on renewable energy sources. After capturing carbon dioxide, it must be sequestered in underground geological formations or deep ocean waters.

Considering safety and environmental concerns for marine life, large capacities of underground porous media, monitoring accessibility, and public acceptance, geological storage seems to be the better option at this time [5]. Underground formations suitable for $CO₂$ sequestration can be either porous media like depleted oil and gas reservoirs and aquifers or non-porous media like host rocks, salt caverns, and abandoned mines. In porous media, gas is stored in interconnected pores or voids previously saturated with saline water or hydrocarbon fluids [6]. Along with depleted reservoirs, saline aquifers are also being explored by the scientific community for large-scale $CO₂$ storage due to their broader availability, versatility with depth and size, reduced leakage and environmental impact (due to mixing with the oil), and potential dual-purpose use as enhanced geothermal systems [7, 8]. There are four primary trapping mechanisms that aid in the safe storage of CO₂ in underground porous media such as structural/stratigraphical, residual, solubility, and mineral trapping [9]. All four trapping mechanisms contribute towards secure $CO₂$ sequestration in the case of subsurface saline aquifers.

An accurate assessment of $CO₂$ -brine displacement dynamics in the pore scale determines the structural, solubility, and residual trapping of $CO₂$. However, limited pore-scale studies have discussed gas displacement dynamics and the corresponding trapping capacity of porous media in terms of both qualitative and quantitative analysis. Moreover, gas-liquid two-phase flow in porous media gets highly complex due to unstable displacement patterns, which are not well investigated. Experimental studies involving microfluidic investigation, 3D X-ray microtomography technique, numerical simulations like lattice Boltzmann method, volume of fluid (VOF) method, and pore-network modeling approach have been utilized by researchers for pore-scale characterization of $CO₂$ trapping and storage capacity estimation [10-17]. Zhang *et al.* conducted a microfluidic investigation and compared the displacement efficiency of nitrogen (N_2) , hydrocarbon gas, and $CO₂$ for miscible and immiscible displacement cases. The discussion focuses on geological $CO₂$ utilization and storage based on enhanced oil recovery and concludes that $CO₂$ is better for oil production, especially in deep reservoirs with high pressure [10]. Al-Zaidi *et al.* and Moghadasi *et al.* used the X-ray microtomography technique to understand $CO₂$ -brine displacement dynamics in core samples in terms of pressure behavior. The unstable displacement dynamics are highlighted in these research articles [11], [14]. Basirat *et al.* used a numerical modeling technique in a 2D microscopic heterogeneous model with the phase field method for interface calculations between $CO₂$ and brine for deep investigation of underground carbon sequestration. The studies show that with a decrease in contact angle, the wetting phase (brine) saturation and normalized interfacial increases [12].

The microfluidic investigation provides a direct visualization during two-phase displacement experiments, allowing for the understanding of unstable displacement patterns like viscous and capillary fingering. In this study, both experimental and simulation techniques have been used to better understand the

displacement dynamics between $CO₂$ and brine. The same microfluidic model was used for both experimental and simulation studies, which mimics subsurface porous media. The study aims to quantify the structural and residual trapping of $CO₂$ and discuss the impact of pore heterogeneity and salinity on storage capacity.

2. MATERIALS AND METHODS

2.1 Numerical Methodology

To simulate $CO₂$ -brine two-phase pore-scale flow, a computational fluid dynamics (CFD) simulation coupling with volume of fluid (VOF) method was used. The VOF model was used to track the interface between the two phases during the simulation. The VOF method introduced a new variable called volume fraction (α), and the corresponding continuity equation for α was solved for the interface tracking. Along with mass, momentum, and energy conservation principles, interfacial forces like viscous and capillary forces were introduced in the model to capture the unstable fingering phenomena at the pore scale. All the governing equations used in this simulation are shown in Table 1.

2.2 Porous Geometry and Experimental Setup

The porous media used for this study has a dimension of 20 mm \times 10 mm \times 20 µm, with 2.5D permeability and 0.57 porosity. For simulation purposes, a two-dimensional (2D) geometry was considered due to the small depth of the three-dimensional micromodel. The average pore size and throat size radii of the physical rock-type microchannel are 46 and 56 μm, respectively. The microfluidic device, as shown in Fig. 1a, has intricate and highly tortuous flow paths and mimics underground porous media. The microfluidic experiments to investigate $CO₂$ -brine displacement dynamics were conducted under a digital microscope with a camera to capture the flow dynamics. Fig. 1b illustrates a simplified experimental setup with a syringe pump for $CO₂/brine$ injection, a microchannel for flow experiments, a digital microscope and camera setup for flow visualization, and a computer for flow investigation. Five different salinities, 0, 10, 30, 60, and 80 parts per thousand (ppt) of water, were used to study the effect of salinity on twophase displacement dynamics. Various salinity values were obtained from seawater (with approximately 30

Table 1 The governing equations used to simulate CO2-brine two-phase immiscible flow in a heterogeneous porous media under subsurface reservoir conditions.

ppt), either by diluting with distilled water or by evaporating the seawater. The water phase was dyed using a blue color dye to differentiate the two phases easily during the flow displacement experiments. Initially, the micromodel was completely saturated with brine. Following the brine saturation, it was left undisturbed for 12–24 hours to achieve a more accurate approximation of subsurface aquifers. $CO₂$ was injected from one side of the microchip to observe the displacement dynamics between the two phases. For each salinity case, experiments were repeated 2-3 times, and average values of flow saturation were reported. After each experiment, the chip was cleaned with ethanol, alkali, and toluene to remove any impurities and previous saline waters.

Fig. 1 (a) Heterogeneous porous media used for the two-phase flow investigation; (b) Experimental setup for microfluidic experiments.

2.3 Pore-scale Studies

All the microfluidic studies were carried out in atmospheric conditions, while numerical simulations were used to investigate under atmospheric and aquifer conditions (P = 15 MPa and T = 323 K). To understand the flow displacement regimes, all the simulation and experimental cases were plotted on a capillary number (*Ca*) vs. viscosity ratio (*M*) phase diagram, as shown in Fig. 2. *Ca* is defined as the ratio of the viscous force of

injected fluid to capillary force ($Ca = \frac{\mu_2 u}{\sigma}$ $\frac{2\mu}{\sigma}$) and *M* is the ratio of viscous force between displacing and displaced fluid ($M = \frac{\mu_2}{\mu_1}$ $\frac{\mu_2}{\mu_1}$). All the considered cases lie between the viscous and capillary fingering regions, which allows to study the influence of viscous and capillary forces on displacement dynamics at the pore scale. The displacement regimes are divided following the works of Lenormand et al. and Zhang et al. [23, 24].

Fig.2 Log (Ca) vs. Log (M) phase diagram illustrating cases considered for this study and three different flow regimes following the work of references [23, 24].

Before carrying out the simulations discussed in this study, the CFD model was checked for mesh independence and validated by comparing it with microfluidic experiments. A detailed discussion on validation is available in our previous literature [20]. Along with atmospheric pressure, simulations were performed under high-pressure and high-temperature conditions observed in subsurface reservoirs. $CO₂$ becomes supercritical under these conditions, and its properties are calculated following the work of Span and Wagner [25]. The drastic change in $CO₂$ density and viscosity observed due to high pressure and temperature cases can be visualized in Fig. 3a and b. Brine and interfacial properties at these conditions were determined following previous literature [20, 26, 27].

3. RESULTS AND DISCUSSIONS

3.1 Experimental Observations

All the microfluidic experiments were conducted under atmospheric conditions to observe the brine displacement patterns and corresponding flow dynamics. The microfluidic device consists of multiple high-permeable, mid-permeable, low-permeable, and dead-end regions to mimic a physical rock found in subsurface saline reservoirs on a quasi-2D scale. A typical observation for a microfluidic experiment is

Fig. 3 Variation of CO² properties at high-pressure and high-temperature geological aquifer conditions (a) Density; (b) Viscosity.

demonstrated in Fig. 4a. Initially, a clean microchip was saturated with brine, and $CO₂$ was flooded. In atmospheric conditions, the interphase is clearly differentiable, indicating an immiscible phase displacement. To understand the physical displacement patterns at the pore scale, 8 locations on the microchip were identified with different permeabilities and at various distances from the injection side. Fig. 4b-d illustrates a few critical observations from the microfluidic experiments. $CO₂$ initially prefers a highpermeable flow path, creating multiple finger-like structures and channeling through these high-permeable channels (Fig. 4b and c). Viscous fingers developed due to the high viscosity contrast (M =0.01) between brine and $CO₂$, which can be observed throughout the micromodel. Fingers developed in the high-permeable

path grow faster and create a positive feedback mechanism, allowing subsequent injected $CO₂$ to follow the same paths [20, 28]. Fingering patterns of mid and low-permeable flow paths only develop after the invasion of high-permeable micro channels (Fig. 4b). Fig. 4c illustrates the CO₂-channeling phenomena and poor sweep efficiency of dead-end regions. This $CO₂$ channeling through high-permeable channels limits the stratigraphical trapping of $CO₂$. The average $CO₂$ saturation observed for 10 ppt salinity brine invasion is 0.495. Fig. 4d shows the snapped-off $CO₂$ bubbles generated as the non-wetting phase separates from the continuous phase and moves separately.

Fig. 4 Microfluidic experiments for the study of CO2 brine (10 ppt) displacement dynamics. (a) A typical CO² flooding observation; (b) Different degrees of viscous fingering; (c) CO² channeling (d) CO² snap-off events.

3.2 Observations from CO² Pore-scale Simulations

CO2-brine two-phase numerical simulations were carried out both at atmospheric and reservoir conditions. Under reservoir conditions, due to the supercritical nature of $CO₂$, breakthrough time and saturation at breakthrough increased significantly. The displacement patterns obtained from CFD simulations highlight the unstable nature of the brine invasion on a pore scale. For 10 ppt salinity values at the atmospheric conditions, Ca and M lie in the intermediate region between capillary and viscous fingering, and the corresponding displacement patterns are demonstrated in Fig. 5. The displacement patterns show unstable fingering patterns and snapped-off $CO₂$ bubbles. Snapped-off bubbles are generated when a gas bubble elongates while traveling through a narrow porous channel and gets disconnected from the continuous phase. A comprehensive understanding of the snap-off effect can help better

estimate CO₂ residual trapping inside a porous media. Higher degrees of snap-off events increase the residual trapping of $CO₂$ and may also increase the solubility trapping as the interfacial area between $CO₂$ -brine increases with each snap-off event. Along with the channel geometry, the degree of snap-off events depends upon the flow rate of the injection, interfacial tension (IFT), and contact angle between rock-brine- $CO₂$. This study observed higher snap-off events for lowsalinity brine, which can be attributed to their low contact angle and low IFT. The $CO₂$ breakthrough time for 10 ppt saline water is 0.739 s, and the CO2 saturation at breakthrough is 0.367. Microfluidic experiments generally estimate higher $CO₂$ saturation as these studies calculate saturation values after a steady state. However, in simulation cases, saturation values were obtained immediately after the breakthrough of $CO₂$ at the outlet. Comparing $CO₂$ saturation at breakthrough can be a better practice to have safe and efficient storage.

time stages of CO2-brine for 10 ppt salinity. Breakthrough time = 0.739 s and saturation at breakthrough = 0.367.

3.3 Impact of Salinity on CO² Trapping

Brine with five different salinity values, 0, 10, 30, 60, and 80 ppt, were used to investigate the impact of brine salinity on $CO₂$ trapping by stratigraphical and residual mechanisms and CO₂-brine displacement patterns. Simulations were conducted at atmospheric and highpressure (P= 15 MPa) and high-temperature (T= 323 K) conditions. Fig. 6 shows a comparative plot of breakthrough time and saturation at breakthrough for different simulation cases. As illustrated with increased brine salinity, both breakthrough time and saturation increase. The trend is similar for both atmospheric and reservoir conditions. After 60 ppt salinity, $CO₂$ saturation remains constant, suggesting the influence of high

viscous force dominating interfacial force compared to other salinity values. Even though the trend is similar for both atmospheric and aquifer conditions, the breakthrough time and $CO₂$ saturation increased by an average of 37.5 % and 29.1 % for underground aquifer conditions. The results show higher $CO₂$ storage capacity of high-temperature and high-pressure reservoirs due to the supercritical nature of $CO₂$. High density and high viscosity at reservoir conditions (Fig. 3) increased the time taken by $CO₂$ to reach the outlet of the micromodel.

Fig. 6 CO² breakthrough time and CO² saturation at breakthrough for different salinity values both at atmospheric and geological reservoir conditions.

CO2-brine displacement dynamics highly depend upon the interfacial forces, IFT, and contact angle. The variation of these two parameters determines the $CO₂$ storage capacity of the micromodel with varying salinity. Fig. 7a and b show the trend of IFT and contact angle for different salinity values at atmospheric and reservoir conditions. Both interfacial tension and contact angle increase with increased brine salinity for atmospheric and reservoir conditions. While in reservoir conditions, IFT decreases drastically, and the contact angle increases, making the system less water-wet. As a result, high-salinity brine is expected to displace easily, allowing more $CO₂$ storage capacity. Similar observations are obtained from the simulation results under extreme reservoir conditions.

4. CONCLUSION

CO² sequestrations will play a crucial role in achieving a net-zero economy. Subsurface saline aquifers have been explored as a potential site for large-scale, safe, and efficient carbon storage. In this context, this study explores the pore-scale flow dynamics of $CO₂$ brine, which helps understand $CO₂$ storage capacity and

Fig. 7 (a) Interfacial tension and (b) Contact angle for different salinity values at atmospheric and reservoir conditions.

leakage status for a porous media filled with saline water. Five different salinity values, 0, 10, 30, 60, and 80 ppt, were considered to evaluate the impact of porescale heterogeneity and salinity on $CO₂$ -brine two-phase flow in a heterogeneous porous microchannel. Microfluidic experiments were performed under atmospheric conditions, and numerical CFD simulations were conducted at atmospheric and high-pressure hightemperature conditions. Microfluidic experiments show the CO2-brine displacement dynamics at various sections of the microchip. Degrees of viscous fingering, $CO₂$ channeling, and snap-off events characterize the $CO₂$ brine displacement dynamics at the pore scale. For reservoir conditions, $CO₂$ storage capacity increased by 29.1 % due to the high density and viscosity of

supercritical $CO₂$. Storage capacity increases with the increase in salinity as the system becomes water-wet with higher salinity. Saline aquifers with a 30-60 ppt salinity can be used to get maximum storage capacity with a balance of viscous and capillary forces. Higher salinity also shifts the displacement dynamics slightly towards the capillary fingering regime. This combined approach of microfluidic experiments and CFD simulations provides a visualization of $CO₂$ -brine twophase displacement dynamics and helps predict the storage capacity of porous media, which would contribute towards a successful large-scale $CO₂$ sequestration.

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